

Significance of tree roots to preferential flow in soil horizons with different degrees of hydromorphy

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- To figure out the influence of roots on the porosity and therefore on infiltrability in hydromorphic soils.
- To assess the significance of the degree of hydromorphy for the relation between roots and hydrologic properties.

2. Methods

- Study site: “Under Scheidwald”, 30 km south of Bern (Switzerland) at an altitude of 1000 m. The forest stand is classified as a **Bazzanio-Abietetum** (Ellenberg & Klötzli, 1972).
- Soil classification: **stagnic gley** (FAO 2000, Waldböden der Schweiz, Band 3, 2006), bulk density between 0.19 and 1.43 g cm⁻³.
- Sprinkler **irrigation** (1m²) close to the stem (16 plots)
- **70 mm of water** application within 1 hour.
- Irrigation **three times** within a 23-hour interval.
- Temporal measurements of soil moisture with **TDR probes** during and after the irrigation.
- Sampling of **root probes** with a soil corer, exactly at the same position where TDR-probes were inserted. Analysis with “WinRHIZO”.
- **The basic units for the investigation were morphological horizons**



1. Degree of hydromorphy
2. Percentage of organic material in non-hydromorphic horizons

Attribute	HG				
	1	2	3	4	5
Topsoil, non-hydromorphic	■				
Cambic		■	■		
Manganese concretions			■	■	
Iron mottles			■	■	
Reduction			■	■	■
Example of horizon*	Ah	B	BGo	Go	Gr

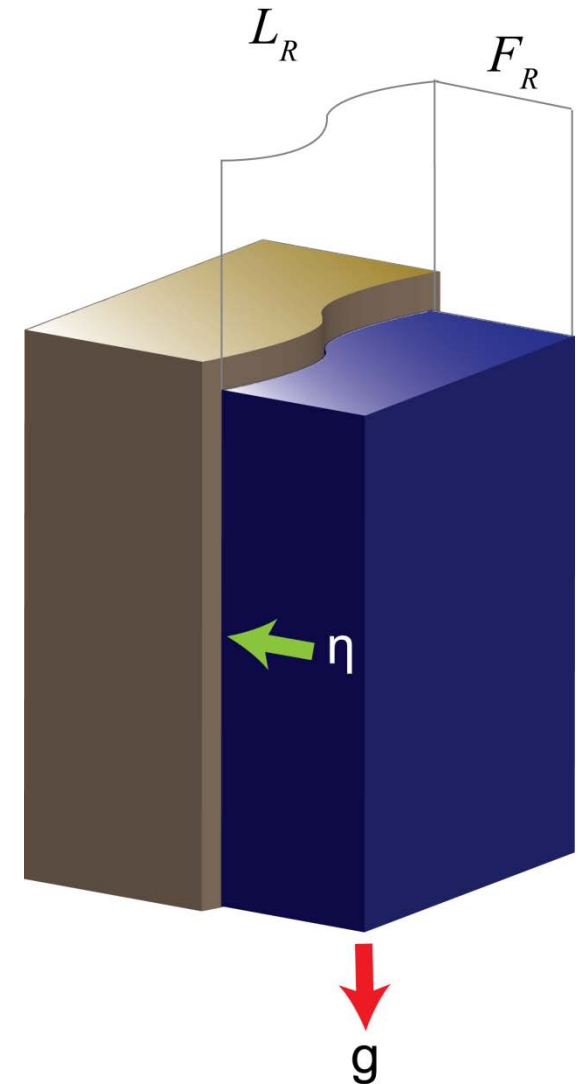
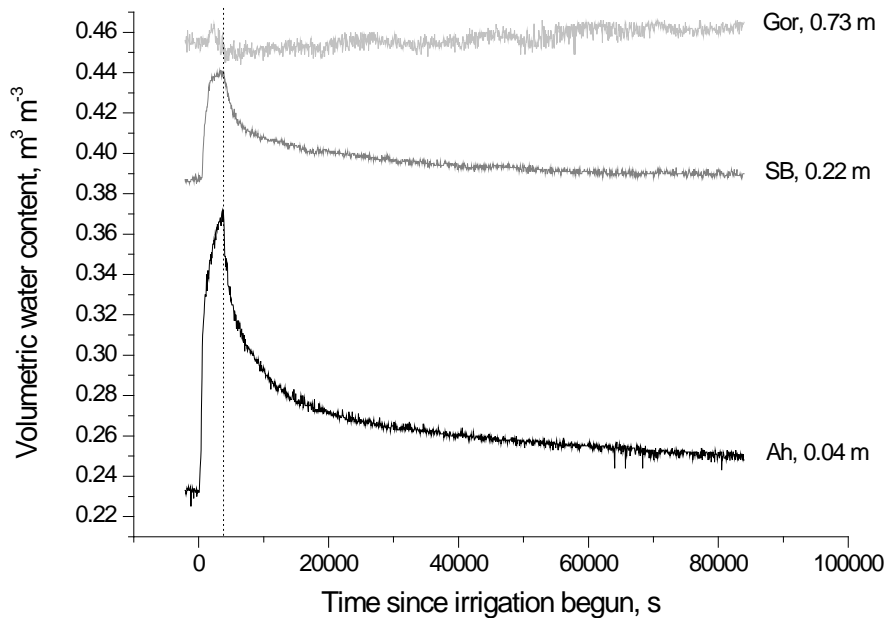
■ Mandatory attribute

■ At least one of the two attributes are mandatory

■ Optional attribute

HG: horizon group. *according to Benzler *et al.* (1982)

- Tiny water streaks (rivulets) are the basic units of preferential infiltration.
- Rivulets are gravity driven and viscosity controlled.
- A rivulet consists of water film that is defined with its thickness F and contact length per area L .



1. Determination of the arrival time of the wetting front t_W .
2. Calculation of the velocity of the wetting front v :

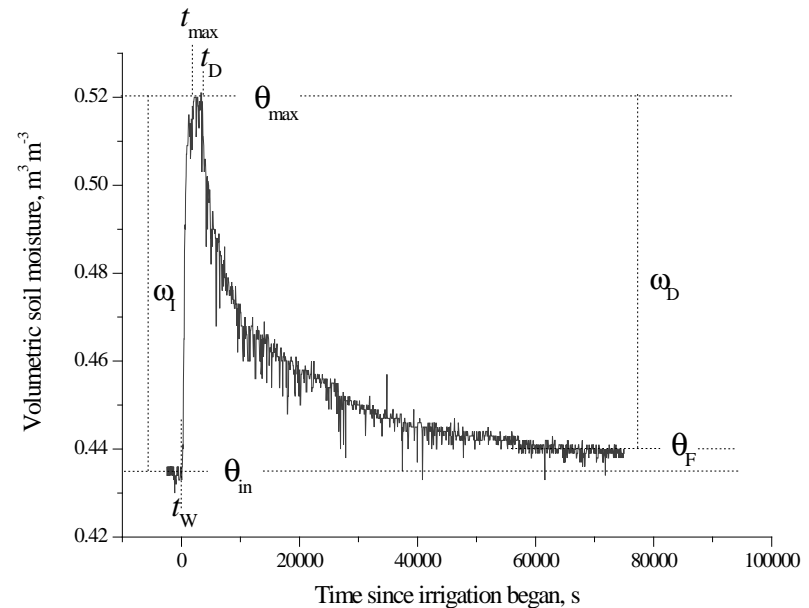
$$v = \frac{Z}{t_W} = \frac{g}{3 \cdot \eta} \cdot F^2$$

3. Calculation of the film thickness F

$$F = \sqrt{\frac{3v\eta}{g}}$$

4. Determination of the water storage capacity w_I
5. Calculation of the contact length L

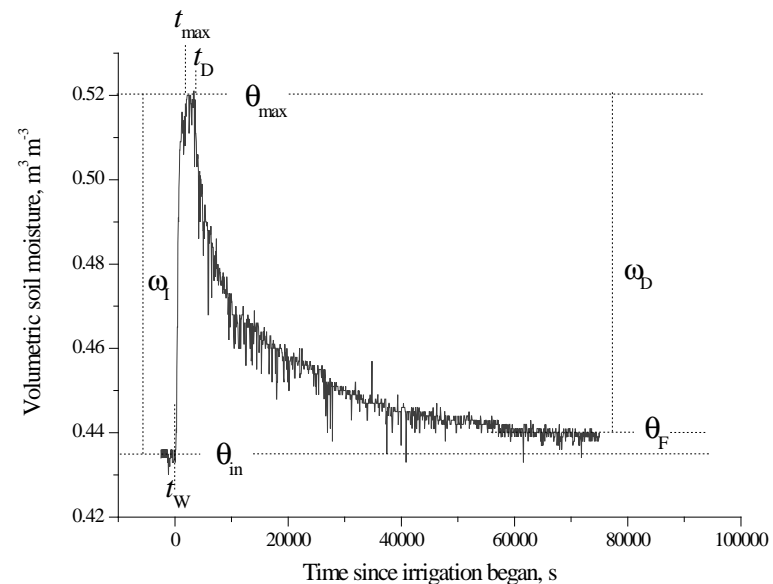
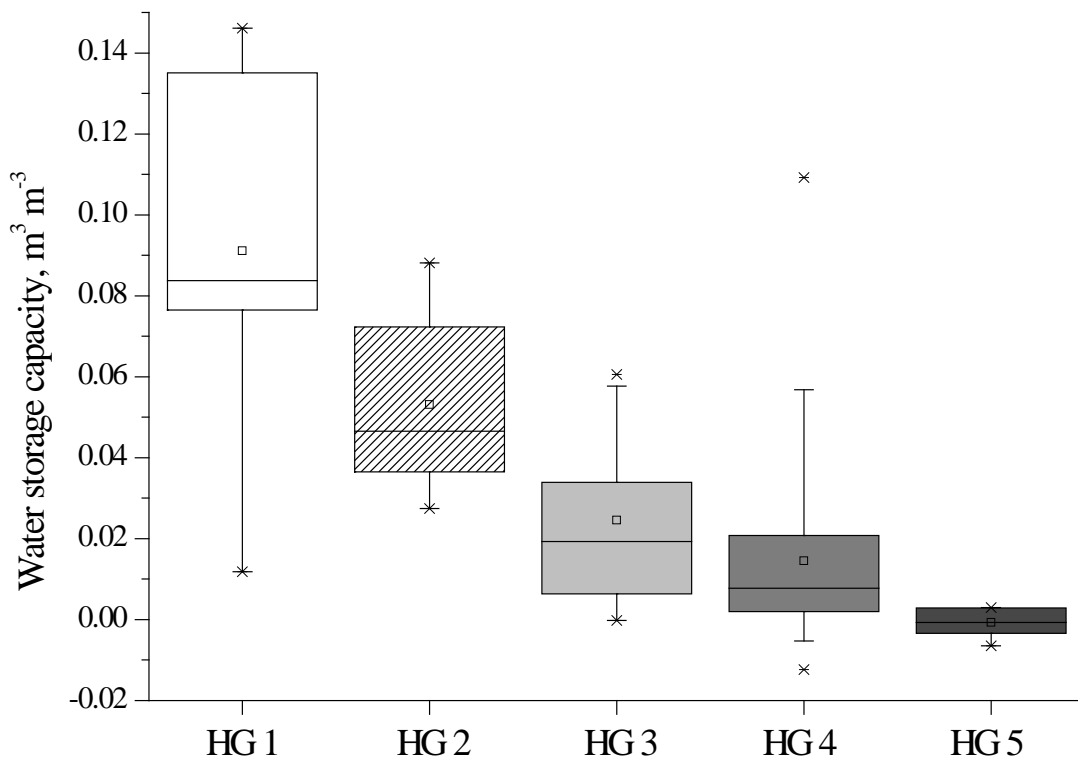
$$w_I = L \cdot F \Rightarrow L = \frac{w_I}{F}$$



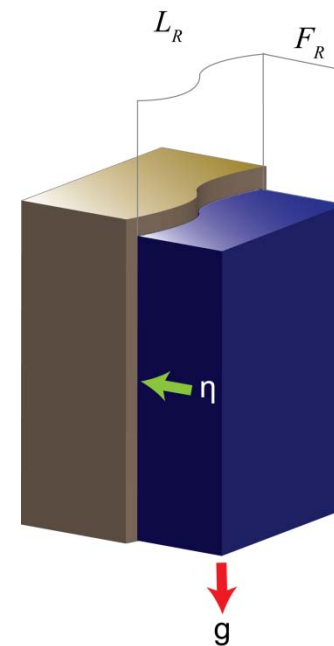
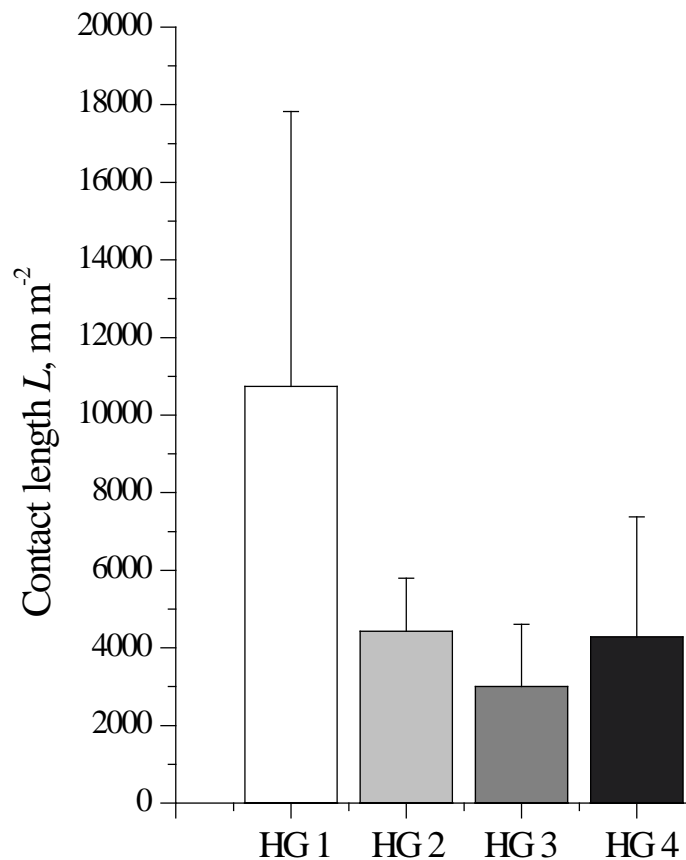
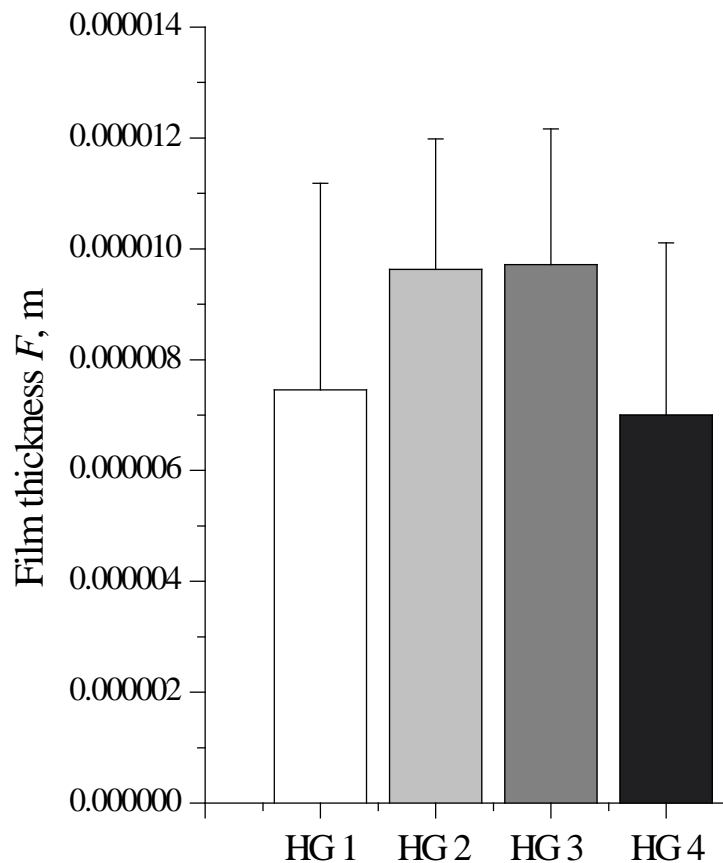
v : velocity of wetting front $m \text{ s}^{-1}$
 Z : depth m
 t_W : arrival time of wetting front s
 g : acceleration due to gravity 9.81 m s^{-2}
 η : viscosity of water $10^{-6} \text{ m}^2 \text{ s}^{-1}$
 F : film thickness m
 L : contact length $m \text{ m}^2$
 w_I : amplitude of infiltration $m^3 \text{ m}^{-3}$

HG	n	First irrigation $m^3 m^{-3}$	Second irrigation $m^3 m^{-3}$	Third irrigation $m^3 m^{-3}$
1	13	0.368 <i>0.137</i>	0.411 <i>0.118</i>	0.415 <i>0.109</i>
2	10	0.359 <i>0.094</i>	0.405 <i>0.066</i>	0.410 <i>0.059</i>
3	20	0.389 <i>0.094</i>	0.414 <i>0.091</i>	0.421 <i>0.085</i>
4	31	0.439 <i>0.066</i>	0.455 <i>0.059</i>	0.454 <i>0.057</i>
5	6	0.440 <i>0.037</i>	0.468 <i>0.031</i>	0.471 <i>0.031</i>

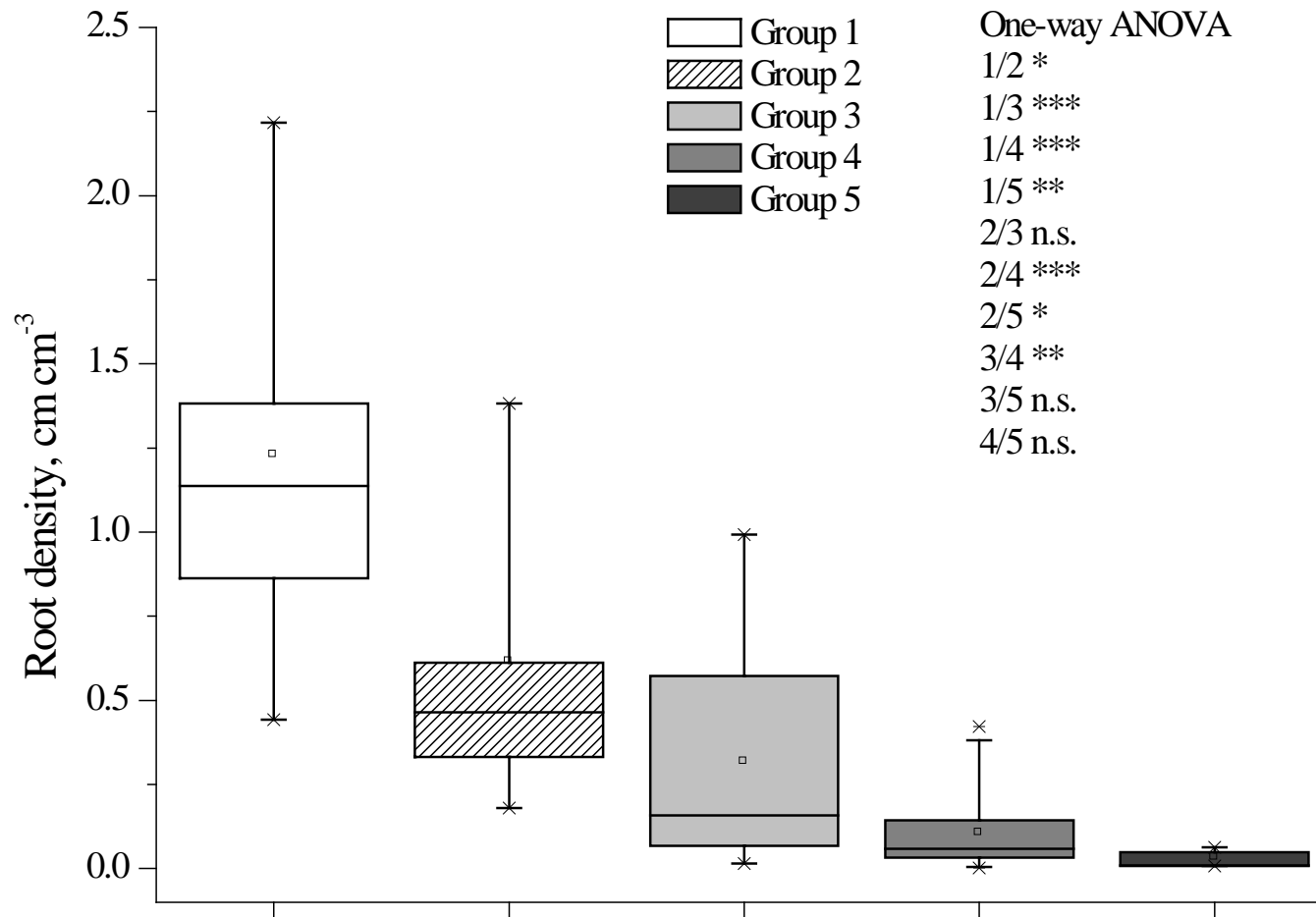
Averaged initial water content of the horizon groups (*italics: standard deviation*): HG 1: non-hydromorphic topsoils; HG 2: non-hydromorphic subsoils; HG 3: few hydromorphic subsoils; HG 4: hydromorphic subsoils; HG 5: waterlogged subsoils; n : sample size



HG 1: non-hydromorphic topsoils; HG 2: non-hydromorphic subsoils; HG 3: few hydromorphic subsoils; HG 4: hydromorphic subsoils; HG 5: waterlogged subsoils



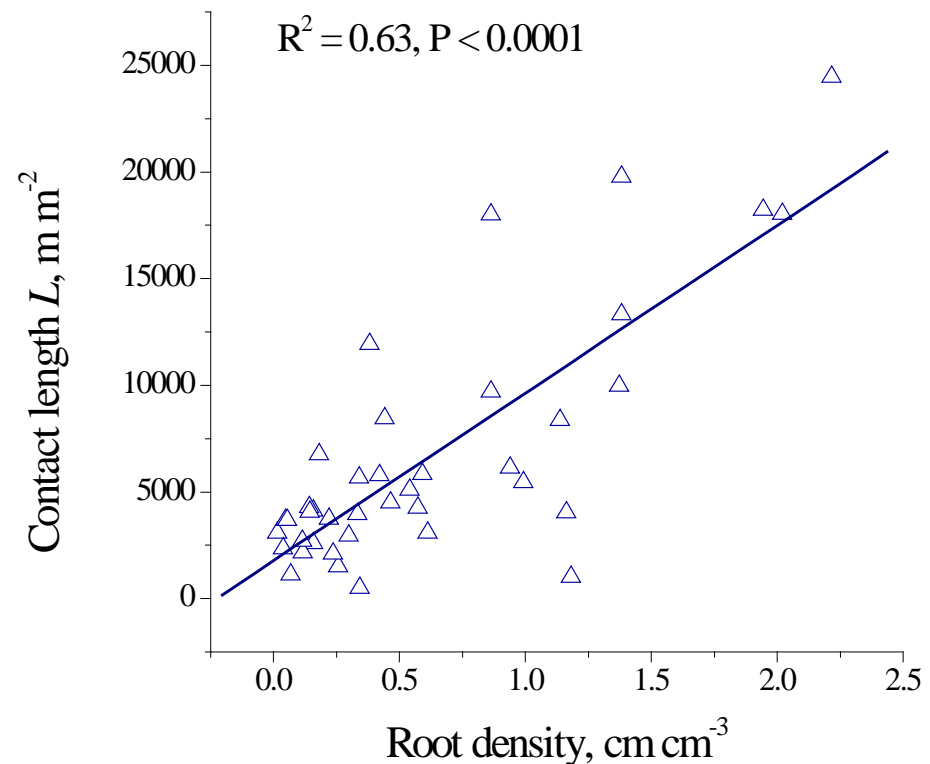
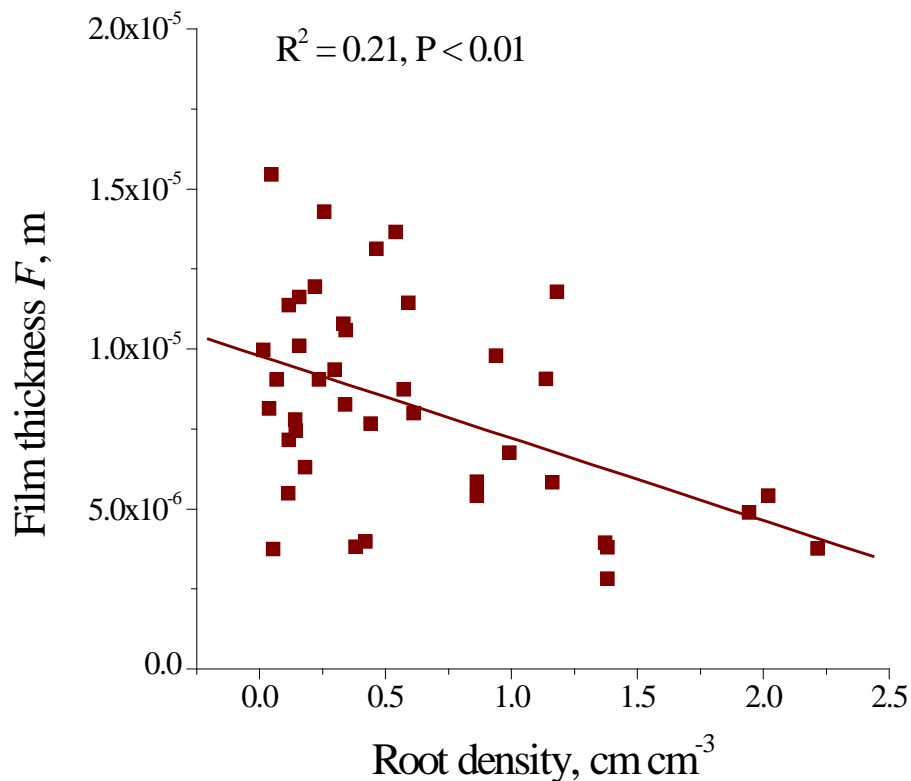
HG 1: non-hydromorphic topsoils; HG 2: non-hydromorphic subsoils; HG 3: few hydromorphic subsoils; HG 4: hydromorphic subsoils



* $P < 0.05$; ** $P < 0.01$; *** $P < 0.001$; n.s.: not significant

Group 1: Non hydromorphic topsoil horizons (Ah)
Group 2: Non hydromorphic subsoil horizons (e.g. B)
Group 3: Few hydromorphic subsoil horizons (e.g. Bcn)
Group 4: Hydromorphic subsoil layers (e.g. Go)
Group 5: Waterlogged subsoil layers (e.g. Gr)

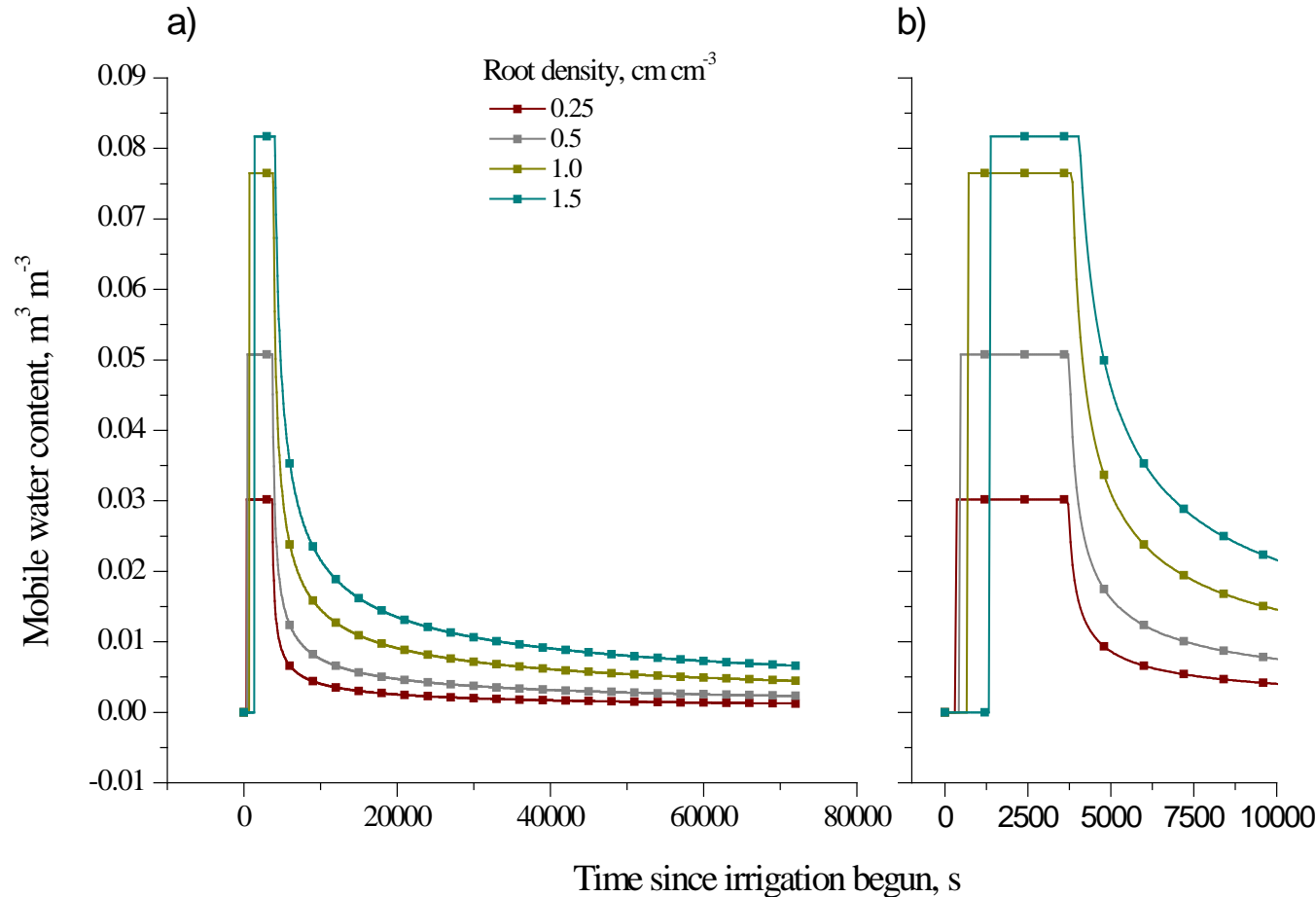
- Application of a multiple linear regression analyses to identify soil properties which determine contact length L .
- Input data: root length density, bulk density, initial water content, pH, texture (percentage of sand, silt and clay)
- Best predictor for L : root length density ($R^2 = 0.63$), only the consideration of the silt content improved the model slightly ($R^2 = 0.66$)



Horizon group	<i>RL</i> vs. <i>L</i>	<i>RL</i> vs. <i>F</i>
1	0.53**	0.48**
2	0.21 n.s.	0.00 n.s.
3	0.26 n.s.	0.29 n.s.
4	0.62**	0.29 n.s.

- Roots influence porosity which is effectively involved in water flow in topsoils and hydromorphic subsoils.
- Infiltratbility in cambic horizons seems not to be influenced by the root density. Swelling of clay?

HG 1: non-hydromorphic topsoils; HG 2: non-hydromorphic subsoils; HG 3: few hydromorphic subsoils; HG 4: hydromorphic subsoils; *RL*: root density; *L*: contact length; *F*: film thickness



- Assumptions: F and L follow from the correlation lines between root density, F and L respectively.
- Soil depth is 0.15 m, intensity of irrigation is 70 mm/h, duration 1 h.

- Increasing root density (up to 1.5 cm cm^{-3}) results in increasing water storage capacity.
- Root densities exceeding 1.5 cm cm^{-3} may decrease water storage capacity, but this process is not relevant since root densities were mostly below 1.5 cm cm^{-3} .



- Root density influenced the pore system accessible to mobile water in topsoils and hydromorphic subsoils when initial water content was high.
- An increase in the root density by 50 % in hydromorphic horizons would increase the water storage capacity from ≈ 0.02 to $0.04 \text{ m}^3 \text{ m}^{-3}$. A 30-cm hydromorphic horizon would be able to store additionally 6 mm water.
- To improve water retention, we propose the cultivation of deep-rooting tree species capable of penetrating into hydromorphic horizons



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